



Trend Analizi Yöntemleri Kullanılarak Doğu Anadolu Bölgesi Aylık Yağış Miktarlarının Değerlendirilmesi

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Özet

Bu çalışmada, Doğu Anadolu Bölgesi'ne düşen aylık ve yıllık toplam yağışların trend analizinin araştırılması hedeflenmiştir. Bu amaçla, Meteoroloji Genel Müdürlüğü'nün 46 adet yağış gözlem istasyonundan alınan, 1960 ile 2013 yılları arasında değişen, verilere Run testi ve Pettitt testi uygulanarak homojenlik analizi yapılmıştır. Homojen olduğu belirlenen istasyonlara Mann – Kendall testi ve Spearman'ın Rho testi uygulanarak trend analizi incelenmiş, Sen'in eğim metodu kullanılarak trendlerin eğimi belirlenmiştir.

Aylık toplam yağışların trend analizi değerlendirildiğinde yaz aylarında genellikle yağışların artan yönde eğilimde olduğu, kış aylarında ise azalan yönde eğilimde olduğu görülmüştür. Yazın Haziran ayında bölgede azalan yönde bir eğilim hâkimken, Temmuz ayında yerini artan yönde bir eğilime bırakmaktadır. Kasım ayı olduğunda ise bölgeye düşen aylık toplam yağışlarda tekrardan azalan yönde eğilim hakim olmaktadır.

Anahtar Kelimeler: Doğu Anadolu Bölgesi, Trend Analizi, Mann–Kendall Testi, Spearman Rho Testi, Sen Eğim Metodu.

Evaluation of Monthly Precipitation Amount in Eastern Anatolia Region by Using Trend Analysis Methods

Abstract

In this study, the research of trend analysis of total monthly and annual rainfalls fall to East Anatolia Region have been targeted. For this purpose, with applying Run test and Pettitt test to data that is taken from General Directorate of Meteorology's 46 units rainfall observation station, changing from 1960 to 2013, homogeneity analysis was performed. With performing Mann-Kendall test and Spearman's Rho test to station determined as homogeneous trend analysis was examined, with using Sen's slope method trend's slope was determined.

As to assessed monthly total rainfalls trend analysis, it is observed that in the summer months generally rainfalls are increasing direction tendency, in the winter months it is decreasing direction tendency. In summer in June when a decreasing direction tendency is dominated at region, in July it gives place to increasing direction tendency. When it is November, again decreasing direction tendency at monthly total rainfalls to region is dominated.

Keywords: East Anatolia Region, Trend Analysis, Mann–Kendall Test, Spearman's Rho Test, Sen Trend Slope Method.

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1. Introduction

Precipitation is the most important component of the complex hydrological cycle with its effect in our daily life. For this reason, precipitation is usually taken as a starting point for understanding the changes in the processes that direct climate. Precipitation data is mostly recorded on land surface and an important element for monitoring the hydrological cycle (Acar and Şenocak 2008).

Long term mean annual rainfall in 631 mm in Turkey while it was determined to reduce by 15% in 1999 and 7% in 2000. In addition to the decrease in the mean annual rainfall, the deviation in the precipitation regime is also a consideration to be emphasized.

In addition, this decrease in rainfall and the deviation in precipitation regime adversely affect agricultural production. If these conditions continue to cause droughts in the coming years, water – related larger problems may occur in the future (Türkeş 1999, Öztürk 2002). However, the sudden decrease in subtropical zone rainfall started to be effective in the eastern Mediterranean Basin and Turkey from 1970s. Significant decreasing trends in rainfall and droughts have emerged as more pronounced in winter. In addition, the Aegean, Mediterranean, Marmara and Southeastern Anatolia regions were the most affected by the arid conditions in the 20-25 years between the early 1970s and the mid-1990s (Türkeş 1998).

There are several studies conducted on the trend analysis of precipitation, flow and temperature data. Demirci et al. (2009) used Thiessen and Mann-Kendall method for the trend of annual, seasonal and monthly mean, maximum and minimum temperatures over the 32-year period between 1975 and 2006 in Istanbul; The tendency of average, maximum and minimum temperatures was determined.

Demirci et al. (2009) used Thiessen and Mann-Kendall method for the annual, seasonal and monthly data obtained over a 32-year period between 1975 and 2006 in Istanbul. The tendency of average, maximum and minimum temperatures was determined. As a result of this study, in the 32-year period for the province of Istanbul, it was observed that the temperatures increased by 0.83°C in annual average temperatures.

Gümüş and Yenigün (2006) investigated the presence and starting point of the trend over the data of 4 current observation stations in the Lower Euphrates Basin using the Mann Kendall method.

Acar and Şenocak (2008) analysed short-term rainfall trends in Turkey's southern province of Adana, which is an important settlement, in time intervals ranging between 5 to 60 minutes. As a result of the study, a significant trend was determined in 5, 10, 15 and 30 minute time scales.

Acar and Senocak (2012) estimated trend analysis with the 1, 6, 12, 24-hour time scale in western Turkey using the 50- year annual data from 7 rainfall stations. After the homogeneity analysis of the data obtained from the rainfall stations during the study, trend analysis was performed with Mann – Kendall test and the slope of the trend was determined by Sen's slope method. As a result of this trend analysis study, a significant trend increase was observed in all time scales for İzmir province, whereas for Antakya and Bursa provinces, a significant trend increase was observed in the 12 - hour time scale. At the same time, breaking year was expressed as 1973 and 1974 for rainfall in the west of Turkey.

Çiflik (2012) investigated the trend of the annual rainfall from 49 DSİ (state water affairs) stations located in the Aegean Region during the 44 - year period between 1962-2005. In this trend study, non-parametric Mann, Kendall, Mann Kendall Mertebe Correlation, Sen-T test and Spearman's rho test were used. According to Mann Kendall, Spearman's rho and T test results, decreasing trend was found in 9 stations.

In the study conducted by Şen (2013), the rainfall and temperature data obtained from the 5 stations in Isparta province were evaluated by using nonparametric Men – Kendall and Spearman's rho test methods. Trend baseline years were evaluated using Mann–Kendall Mertebe Correlation test and trends tendency was determined using Sen's trend tendency method and Linear regression method.

In the study conducted by Uçgun (2010), the trend analysis of rainfall, temperature, evaporation and flow data were performed with the data collected from the observation stations in Kızılırmak Basin. In the analysis of the flow data, decreasing trend was observed in 4 stations.

In the study of Büyükyıldız (2004), the trend analysis of the rainfall falling on the Sakarya Basin was examined and at the same time stochastic modeling was performed. In this trend study, Sen's T test, Spearman's rho test, Mann - Kendall test and seasonal Mann Kendall test were used under two headings, monthly average annual bases and monthly change of rainfall of each station.

Partal and Kahya (2006) determined the significance of monthly and annual precipitation data in Turkey using trend analysis. This study adopted non-parametric methods using the Mann - Kendall test and Sen's T-test. A 64 – year data set between 1929 and 1993 was used over 96 stations.

Güventürk (2013) aimed to determine the rainfall trends in the Euphrates, Tigris, Aras and Çoruh basins located in eastern Turkey. In these four basins, the trend analysis was conducted from 15 stations and evaluated for a 40-year period between 1970 and 2010. As a result of the trend of rainfall, a significant increase was observed in November, December, January and February. In March, April, May and June, there was a slight decrease in large river flows. Seven among 15 selected stations tended to decrease in the number of wet days.

Norrrant and Douguédroit (2006) performed trend analysis on monthly, seasonal and yearly time scales of the Mediterranean geography between the years of 1950 and 2000. Mann - Kendall test method was used over the data obtained from 63 stations in the study.

Mondal et al. (2012) conducted the trend analysis of rainfall in the Orissa River Basin in India. Mann Kendall test and Sen's T Test method were used among nonparametric methods. As a result of the study conducted a 40-year period in the Orissa Basin between 1971 and 2010, it was stated that these trends are not significant despite the determination of both trend slope with increasing and decreasing direction with Mann - Kendall test.

Liu et al. (2012) collected the data from 186 rainfall measurement stations covering the years 1956 to 2000 in the city of Guangdong in southern China and evaluated them using Mann - Kendall test method on monthly, seasonal and annual bases.

Gautam and Acharya (2012) is different from other studies in that study samples in others are evaluated according to the amount of precipitation this study evaluates the trend analysis of flow quantities.

Serrano et al. (1999), conducted trend analysis of the total monthly and annual rainfall in the Iber Peninsula.

In the study conducted by Siddik and Rahman (2014), the data obtained from 15 observation stations in Bangladesh between 1961-2008 were analyzed by using Mann Kendall test method for maximum, minimum and average temperatures. At the same time, Sen's slope method was used to determine the magnitude of the trend.

In the study conducted by Talaei (2014), Mann - Kendall and Spearman's rho test methods were used to analyse the trend of the data from 7 rainfall measurement stations in the west of Iran between 1969 and 2009.

Raziei et al. (2014) performed trend analysis over the data obtained from the Global Precipitation Climatology Center in Iran for a 50-year period between 1951 and 2009 considering monthly, seasonal and annual rainfall values.

Han et al. (2013) completed a trend analysis over river flows and precipitation trends in the Xiangxi River Basin on monthly, seasonal and annual bases using the Mann - Kendall test and Sen's slope method.

In some recent studies, it has been pointed out that the annual rainfall in the Mediterranean and Aegean Regions, the decrease in the total rainfall in the winter season is strong (Demir et al. 2007, Türkeş et al. 2007).

It may be helpful to know the change in the amount of water over time and use, store and plan more carefully water. For this reason, there is a significant need for statistical methods such as trend analysis (Gümüş and Yenigün 2006). The better the trend analysis of rainfall is, the better to know and determine the future of water. A study covering the Eastern Anatolia Region was conducted in order to determine the past and the future of the rainfall. With this study, it is aimed to determine the trends in the last 50 years on a monthly basis with the monthly rainfall data obtained from the stations within the borders of the Eastern Anatolia Region of the State Meteorological Directorate. In this study, in Eastern Anatolia covering 163,000 km², 21% of Turkey, non-parametric methods of Mann-Kendall test and were used to determine rainfall trend and to Sen's slope test to determine the trend slope.

2. Materials and Methods

2.1. Study area and Data

East Anatolia Region is the east of Turkey and roughly of the triangular shape and the most elevated part of the country. The region is bordered by the south faces of the North Anatolian Mountains in the north and the north high skirts of the Taurus Mountains in the south. Regional boundary separating the region from the Interior Anatolia follows the rough water section of the line of the Kızılırmak and Euphrates rivers and with 163,000 square kilometers of surface area, the region constitutes 21% of the Turkey's surface area, which carries the region to the first place among seven geographic regions.

The main surface characteristics of the region are the folded mountain ranges and the plains and plateaus between them. The average elevation of the major part of the region does not fall below 1500 meters, and more than half of the region's surface area is located at an altitude of 2000-2500 m.

East Anatolia Region is deeply inserted into the continental - originated climatic regions of Asia and surrounded by high mountain ranges from north and south. For this reason, the region is closed to the favourable climatic effects of the Mediterranean and Black Sea. Although it differs within the region, the semi-arid, arid continental climatic characteristics with long, severe and hard winters are dominant in the region. Depending on the altitude, the average annual temperature is between 3 and 13°C. The most severe winter conditions are experienced especially in the northeastern part of the region. Due to lower humidity content in the region where terrestrial dominance is observed, daily temperature differences are high and evaporation rate is also high in the southern parts. The region is rich in water resources.



Figure 2.1. Map including provinces in East Anatolia

2.1.1. Data

Monthly rainfall data to be used in the study were collected from 46 meteorological rainfall measurement stations of the General Directorate of State Meteorological Affairs (Table 2.1).

Table 2.1. Rainfall measurement stations considered in the study

Station	Longitude (⁰ E)	Latitude (⁰ N)	Elevation (m)	Record Duration
Ağrı	43.05	39.72	1646	1960-2013 (54)
Doğubeyazıt	44.08	39.55	1725	1963-2009 (47)
Ardahan	42.70	41.10	1827	1961-2013 (53)
Erzincan	39.48	39.75	1216	1960-2013 (54)
Tercan	40.39	39.77	1429	1964-2013 (50)
Erzurum	41.18	39.95	1758	1960-2008 (49)
Hınıs	41.69	39.36	1715	1964-2001 (48)
Horasan	42.17	40.03	1540	1969-2013 (45)
İspir	40.99	40.48	1223	1965-2011 (47)
Oltu	41.99	40.54	1312	1966-2013 (48)
Tortum	41.54	40.30	1576	1963-2013 (51)
Iğdır	44.05	39.92	856	1960-2013 (54)
Kars	43.10	40.60	1777	1960-2013 (54)
Arpaçay	43.32	40.84	1688	1970-2013 (44)
Sarıkamış	42.59	43.32	2092	1963-2013 (51)
Bingöl	40.50	38.88	1177	1963-2013 (51)
Genç	40.55	38.73	1250	1980-2013 (34)
Solhan	41.05	38.95	1366	1965-2013 (49)
Elazığ	39.25	38.64	989	1960-2013 (54)
Ağın	38.71	38.94	900	1979-2013 (35)
Baskil	38.83	38.57	1300	1979-2013 (35)
Karakoçan	40.04	38.94	1090	1980-2013 (34)
Keban	38.74	38.79	808	1963-2013 (51)
Maden	39.40	38.23	1100	1980-2010 (31)
Palu	39.92	38.69	869	1966-2013 (48)
Sivrice	39.31	38.45	1240	1979-2013 (35)
Malatya	38.21	38.33	950	1960-2013 (54)
Arapgir	38.48	39.04	1211	1964-2010 (47)
Doğanşehir	37.88	38.09	1223	1965-2013 (49)
Tunceli	39.54	39.10	981	1964-2013 (50)
Çemişgezek	38.90	39.05	953	1969-2013 (45)
Mazgirt	39.60	39.01	1400	1982-2013 (32)
Bitlis	42.16	38.47	1785	1965-2010 (46)
Ahlat	42.47	38.74	1730	1968-2013 (46)
Tatvan	42.28	38.50	1687	1965-2013 (49)
Hakkari	43.73	37.57	1727	1963-2013 (51)
Yüksekova	44.28	37.57	1877	1969-2011 (43)
Muş	41.50	38.75	1322	1964-2013 (50)
Malazgirt	42.53	39.13	1540	1963-2013 (51)
Varto	41.44	39.17	1510	1977-2010 (34)
Van	43.34	38.46	1675	1960-2013 (54)
Başkale	44.01	38.04	2286	1964-2013 (50)
Erciş	43.33	39.01	1678	1965-2013 (49)
Gevaş	43.11	38.29	1694	1982-2013 (32)
Özalp	43.97	38.65	2000	1969-2013 (45)
Muradiye	43.76	38.98	1706	1964-2013 (50)

Since the data were not homogeneous at a total of 8 rainfall measurement stations of Ahlat, Ardahan, Arpaçay, Hakkari, Kars, Özalp, Malatya and Yüksekova among 46 ones, these were excluded from the study.

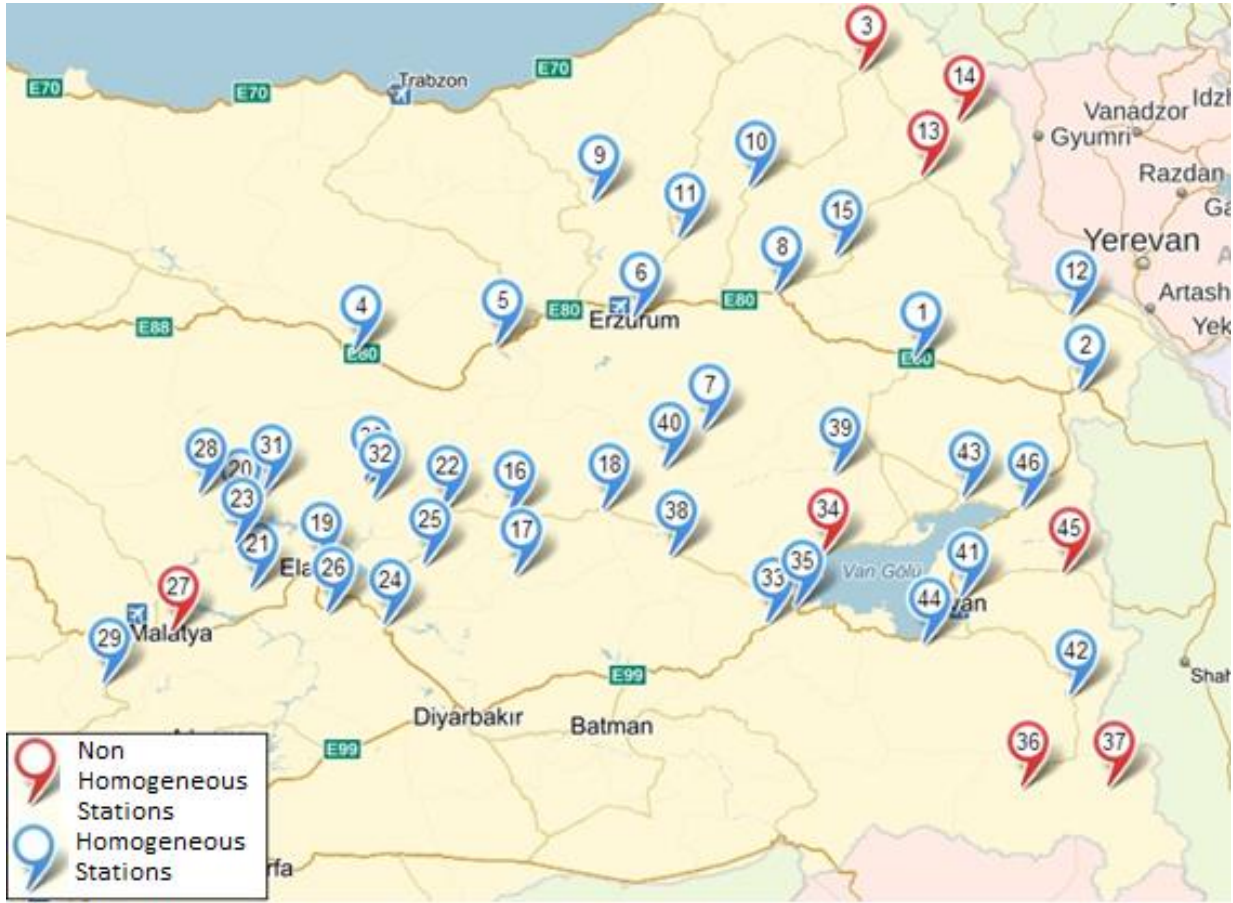


Figure 2.1. Locations of totally 46 rainfall measurement stations.

2.2. Homogeneity

The methods used to determine the trend of a series are collected under two headings; parametric and non-parametric methods, depending on whether the series is tied to a distribution or not (Hirsch 1992).

In parametrical methods, the real value of the data in the series is important and this value is used in the calculations. It is known that non-parametric methods give effective results compared to parametrical methods without the necessity of complying with normal distribution (Hirsch 1992). The Run (Swed Sw Eisenhart) and Pettitt test methods were preferred to determine the homogeneity of the series.

2.2.1. Swed- Eisenhart run test

The Swed and Eisenhart (Swed and Eisenhart 1943) run test is a non-parametric procedure used in determining the homogeneity of a time series. In general, it can adequately detect non-homogeneities as well as other problems such as a change of instrument, relocation of stations, etc. However, after detection of the non-homogeneity, further detailed physical and/or meteorological studies are needed in order to identify its major cause.

The test procedure depends on the truncation of the analyzed time series at the median level, giving rise to data values greater or smaller than the median. Any uninterrupted sequence of greater (or smaller) values preceded and succeeded by at least one smaller (or greater) value is referred to as a run. A succession of greater (or smaller) values constitutes a positive (or negative) run. In general, the number of positive runs is equal to negative runs plus or minus 1. For the analyzed time series to be homogeneous, the number of positive (or negative) runs should be confined within the upper and lower confidence limits at a given significance level (usually 5 or 10 percent). The calculation of confidence limits is based on the assumption of a normal distribution. According to this test, if the number of runs falls between confidence limits than the data set considered is homogeneous (Thorn 1966).

2.2.2. Pettitt test

The Pettitt Test is a nonparametric test that does not require any assumptions about the distribution of data. This nonparametric method developed by Pettitt (1979) to determine the point of change in a time series can find the point of change on a monthly or annual scales. While the null hypothesis indicates an independent and random distribution of the series, the alternative hypothesis indicates a sudden change. The test statistic is associated with the Mann - Whitney statistic. The critical values of this test are given below.

$Y_1, Y_2, Y_3, \dots, Y_n$ values are sequenced as $r_1, r_2, r_3, \dots, r_n$

$$X_k = 2 \sum_{i=1}^k r_1 - k(n + 1) \quad k = 1, 2, \dots, n \quad (2.1)$$

X_k values are drawn as graphics. The absolute maximum value of X_k determines the point of change.

$$X_E = \max_{1 \leq k \leq n} |X_k| \quad (2.2)$$

2.3. Trend

The trend that can be defined as long-term movement is the development or tendency of a time series in a long-term and a certain direction. Since trend analysis is a long-term analysis, the fact that the data is given monthly or seasonal does not affect the result of the analysis. The trend of a series can be linear or curvilinear. However, an important feature of the trend is stable in both cases (Köksal 1998).

2.3.1. Mann–Kendall test

Mann-Kendall test is a statistical testing method that is commonly used in defining the trend in time series in hydrology and climatology fields and that is also suggested by World Meteorological Organization (WMO). Mann-Kendall test pulled its rank in trend analysis in several studies. Non-parametric tests are usually more robust compared with parametric ones, among which the Mann–Kendall test is the most used in hydrology and climatology. Therefore, to detect any monotonic trends in the precipitation time series at all considered grid points, the Mann–Kendall test was used. This test consists of comparing each value of the time series with the remaining in a sequential order. The Mann-Kendall statistic S is given as:

$$S = \sum_{k=1}^{n-1} \sum_{j=k+1}^n \text{sgn}(x_j - x_k) \quad (2.3)$$

The application of trend test is done to a time series x_i that is ranked from $i=1, \dots, n-1$ and x_j , which is ranked from $j=i+1, \dots, n$. Each of the data point x_i is taken as a reference point which is compared with the rest of the data points x_j so that,

$$\text{sgn}(x_i - x_j) = \begin{cases} +1 & \text{if } (x_i - x_j) > 0 \\ 0 & \text{if } (x_i - x_j) = 0 \\ -1 & \text{if } (x_i - x_j) < 0 \end{cases} \quad (2.4)$$

It has been documented that when $n \geq 8$, the statistic S is approximately normally distributed with the mean. $E(S) = 0$

The variance statistic is given as:

$$\text{Var}(S) = \frac{n \cdot (n - 1) \cdot (2n + 5)}{18} \quad (2.5)$$

The test statistics Z is computed as:

$$Z = \begin{cases} \frac{S - 1}{\sqrt{\text{Var}(S)}}, & S > 0 \\ 0 & S = 0 \\ \frac{S + 1}{\sqrt{\text{Var}(S)}}, & S < 0 \end{cases} \quad (2.6)$$

Z here follows a standard normal distribution. A positive (negative) value of Z signifies an upward (downward) trend. A significance level α is also utilized for testing either an upward or downward monotone trend (a two-tailed test). If Z appears greater than $Z(\alpha/2)$ where α depicts the significance level, then the trend is considered as significant.

2.3.2. Spearman's Rho test

A quick and simple test to determine whether correlation exists between two classifications of the same series of observations is the Spearman's rank correlations test. In this test, there is a significant trend only if the correlation between time steps and rainfall observations are found to be significant. Account of the test statistic Z based on r_s was not presented here, since it can easily be found in statistical books. The test statistics r_s is computed as:

$$r_s = 1 - 6 \left[\sum_{i=1}^n (R(x_i) - i)^2 \right] / (n^3 - n) \quad (2.7)$$

For $n > 30$, the distribution of r_s will be normal, so that the normal distribution tables can be used. In this case, the test statistic r_s is computed by

$$Z = r_s \sqrt{n - 1} \quad (2.8)$$

If $|Z| > Z\alpha$ at a significance level of α ; then the null hypothesis of no trend (on the other word, values of observations are identically distributed) is rejected.

2.3.3. Sen's estimator of slope

If a linear trend is present, the true slope (change per unit time) can be estimated by using a simple non-parametric procedure developed. In computational procedures, the slope estimates of N pairs of data are first computed by

$$Q_i = \frac{x_j - x_k}{j - k} \quad (2.9)$$

for $i=1, \dots, N$; where x_j and x_k are data values at times j and k ($j > k$); respectively. The median of these N values of Q_i is Sen's estimator of slope. If there is only one datum in each time period, then

$$N = n(n - 1) / 2 \quad (2.10)$$

where n is the number of time periods. If N is odd, then Sen's estimator is computed by

$$Q_{median} = Q_{(N+1)/2} \quad (2.11)$$

and if is even, then Sen's estimator is computed by

$$Q_{median} = \left[Q_{N/2} + Q_{(N+2)/2} \right] / 2 \quad (2.12)$$

The detected value of Q_{median} is tested by a two-sided test at the $100(1-\alpha)\%$ confidence interval and true slope may be obtained by the non-parametric test.

3. Results

Missing monthly rainfall data obtained from the General Directorate of Meteorology were estimated and homogeneity of all data together with missing ones were analysed for their homogeneity.

The trend analyses and slopes of the data were determined being obtained from the stations determined to be homogeneous. In the evaluations, monthly rainfall data in Eastern Anatolia Region were taken into consideration using non-parametric methods, Mann-Kendall test and Spearman's rho test individually and together. Trend slopes of the stations were determined through Sen's slope method.

The findings obtained were explained on monthly and station bases.

3.1. Homogeneity analysis

As the result of the homogeneity analysis performed by both methods, totally 8 rainfall measurement stations were excluded from the study; one (Hakkari Rainfall Measurement Station) through the Run (Swed - Eisenhart) Test and seven through the Pettitt Test namely the Rainfall Measurement Stations of Ahlat, Ardahan, Arpaçay, Kars, Malatya, Özalp and Yüksekova since the data from these stations were not homogenous. Thus, data from 38 rainfall observation stations were evaluated in the trend study.

3.2. Monthly Trend Analysis

While performing monthly trend analysis, the monthly total rainfall data collected from the stations is analyzed in a way that will be evaluated within the month. Monthly precipitation data corresponds to the same number of data used in the trend analysis of the annual total precipitation data. Monthly rainfall data corresponds to the same number of data used in the trend analysis of the annual total precipitation data. It starts and ends in the same year. In the monthly trend analysis, Mann - Kendall test and Spearman's rho test were used with 90% and 95% confidence interval.

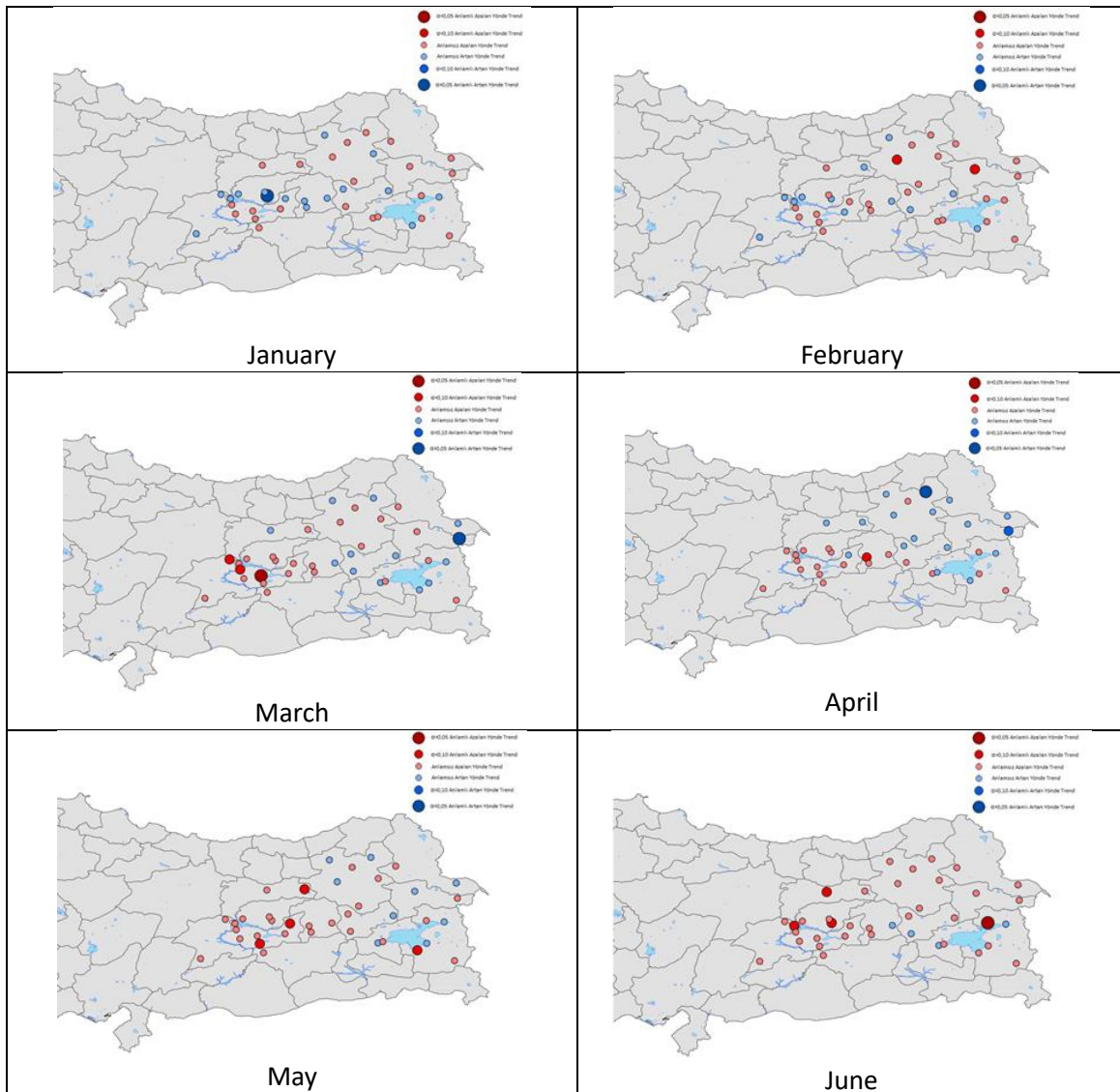
3.2.1. Trend analysis of monthly rainfall with Mann-Kendall test

The results of the analysis conducted through Mann-Kendall test over monthly rainfall at 95% and 90% confidence intervals are given in Table 3.1. The results are also shown in Figure 3.1.

Table 3.1. Mann-Kendall test results for 10% and 5% level of significance

Month	Increasing trend	Decreasing Trend
January	Mazgirt*	-
February	-	Ağrı, Erzurum
March	Doğubeyazıt*	Arapgir, Elazığ*, Keban
April	Doğubeyazıt, Oltu*	Bingöl
May	-	Gevaş, Karakoçan, Sivrice, Tercan ⁺
June	-	Ağın, Erciş*, Erzincan, Mazgirt ⁺
July	Başkale*, Elazığ, Erciş, İspir*, Oltu	-
August	Hınıs, Muş	-
September	Ağın*, Çemişgezek*, Tunceli, Van, Varto*	-
October	-	Erciş, Gevaş
November	-	Ağın, Genç*, Tunceli*, Solhan ⁺
December	-	İspir, Palu*, Solhan*

* Significant at both 5 and 10 per cent levels , + Significant at only 5 per cent levels



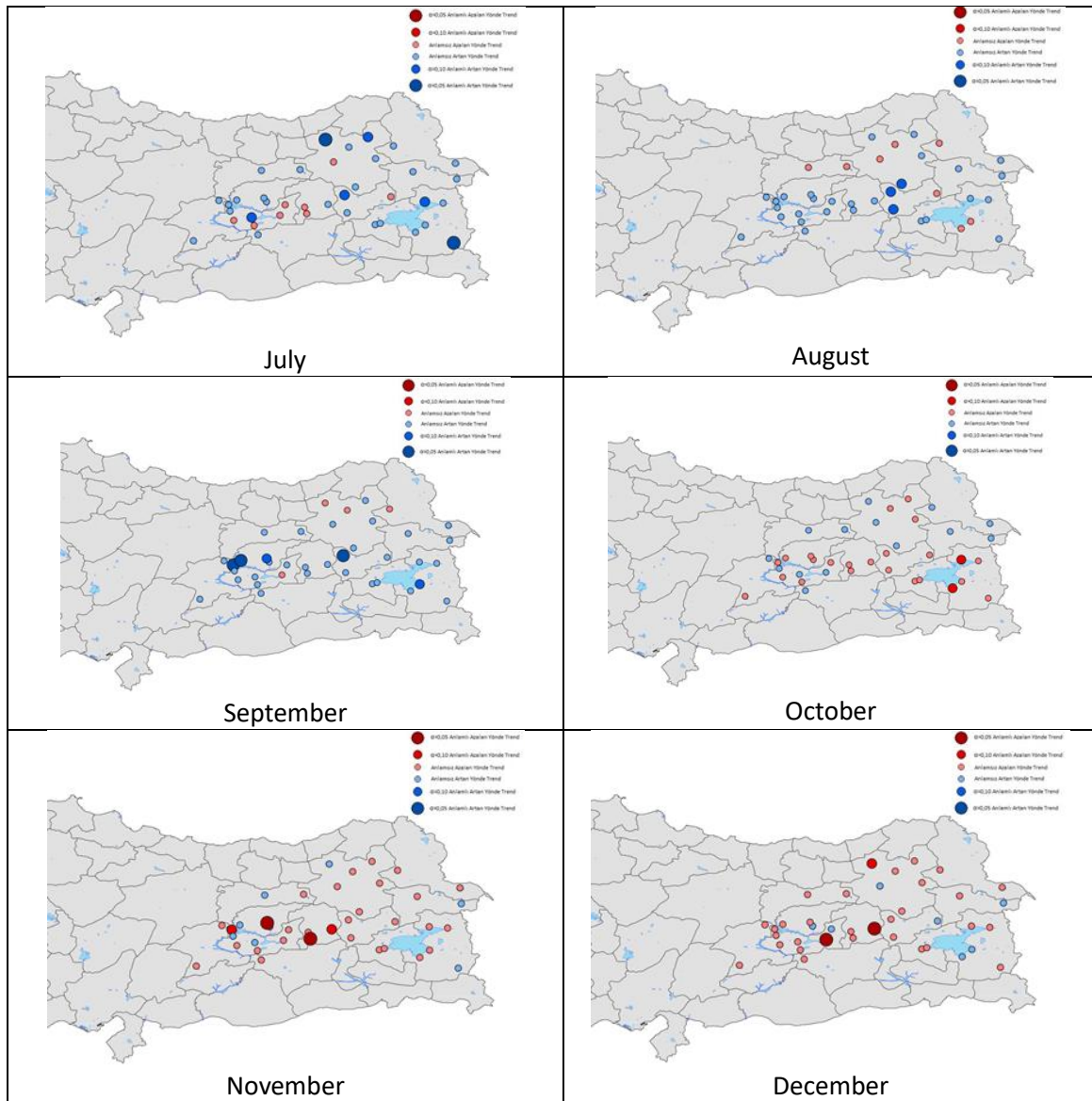


Figure 3.1. Mann-Kendall test results for 10% and 5% level of significance

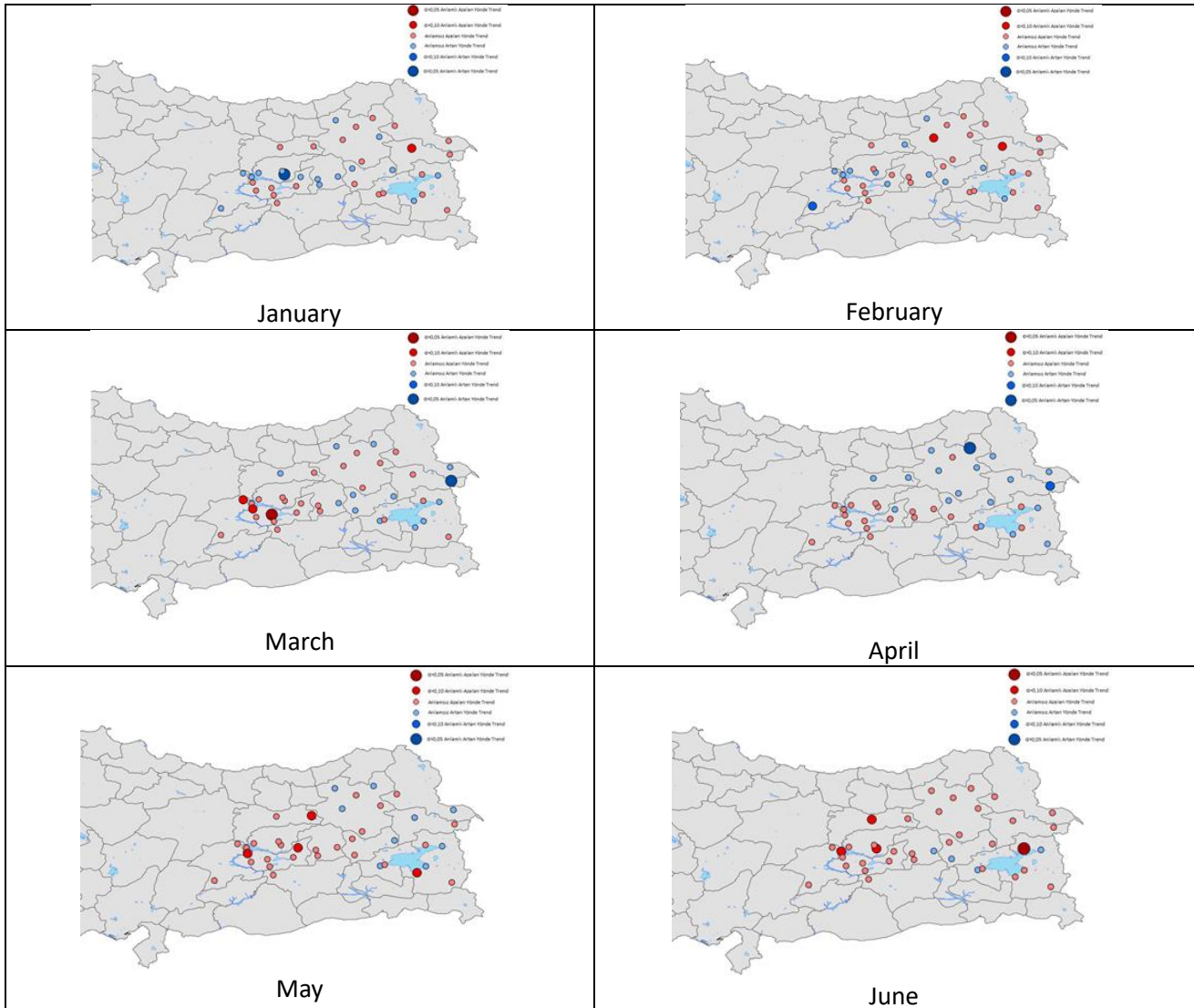
3.2.2. Trend analysis of monthly rainfall with Spearman's rho test

Spearman's rho test method was used to perform trend analysis of monthly rainfall at 95% and 90% confidence intervals are given in Table 3.2. The results are also shown in Figure 3.2.

Table 3.2. Spearman's rho test results for 10% and 5% level of significance

Month	Increasing trend	Decreasing Trend
January	Mazgirt*	Ağrı
February	Doğanşehir	Ağrı, Erzurum
March	Doğubeyazıt*	Arapgir, Elazığ*, Keban
April	Doğubeyazıt, Oltu*	-
May	-	Gevaş, Karakoçan, Keban
June	-	Ağın, Erciş*, Erzincan, Mazgirt
July	Başkale*, Çemişgezek, Doğanşehir* Elazığ*, Erciş*, İspir*, Keban*, Maden, Oltu, Tatvan, Tunceli*, Varto	-
August	Ağın*, Baskil, Bingöl*, Çemişgezek*, Doğanşehir*, Elazığ*, Genç, Hınıs, Karakoçan*, Keban*, Maden*, Mazgirt*, Muş*, Palu*, Sivrice*, Tatvan, Tunceli*, Varto*	-
September	Ağın*, Baskil*, Çemişgezek*, Horasan, Maden, Muradiye, Tunceli, Van*, Varto*	-
October	-	Erciş
November	-	Ağın, Genç*, Maden, Solhan, Tatvan, Tunceli
December	-	İspir, Palu*, Solhan*, Tercan

* Significant at both 5 and 10 per cent levels , + Significant at only 5 per cent levels



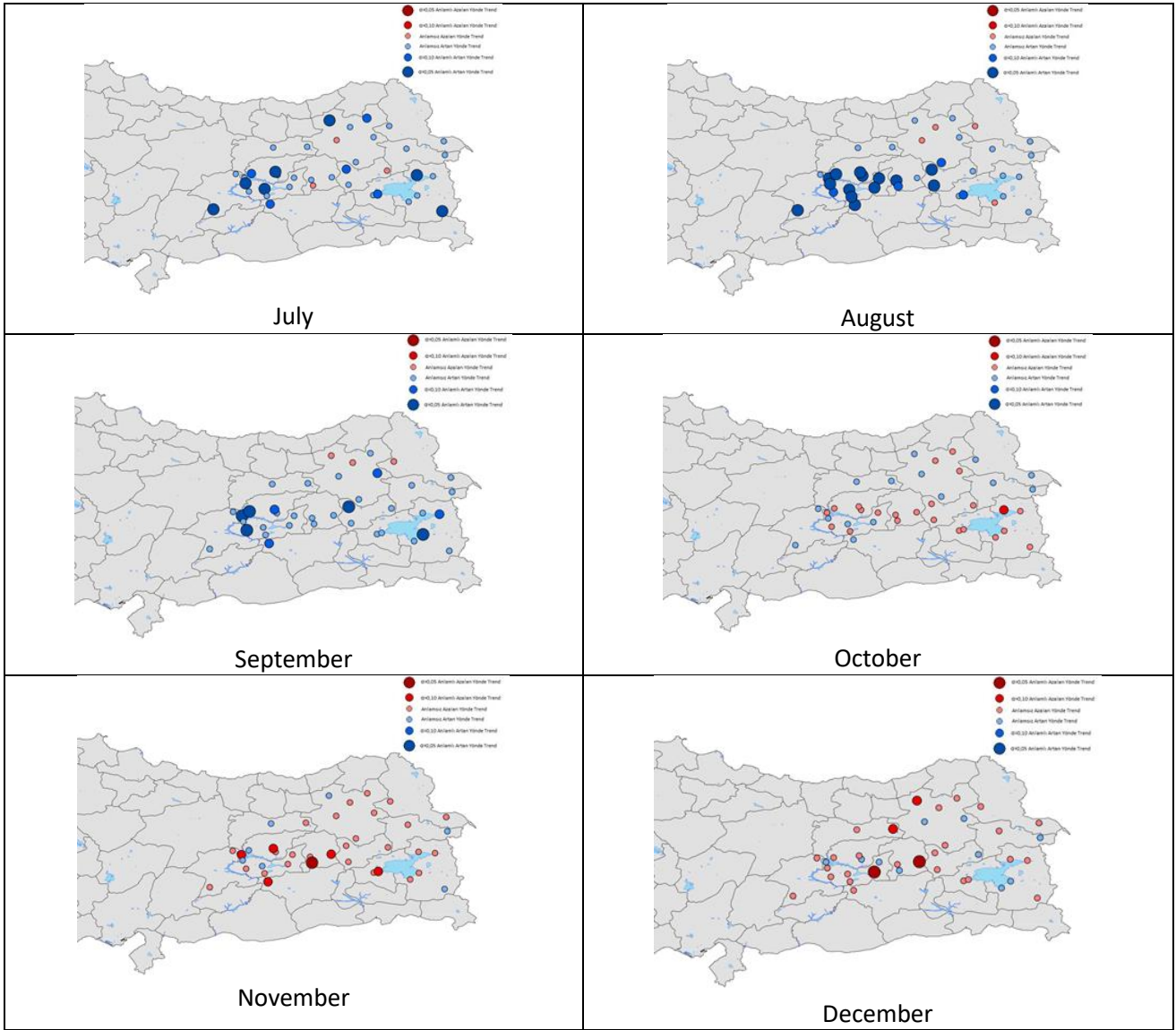


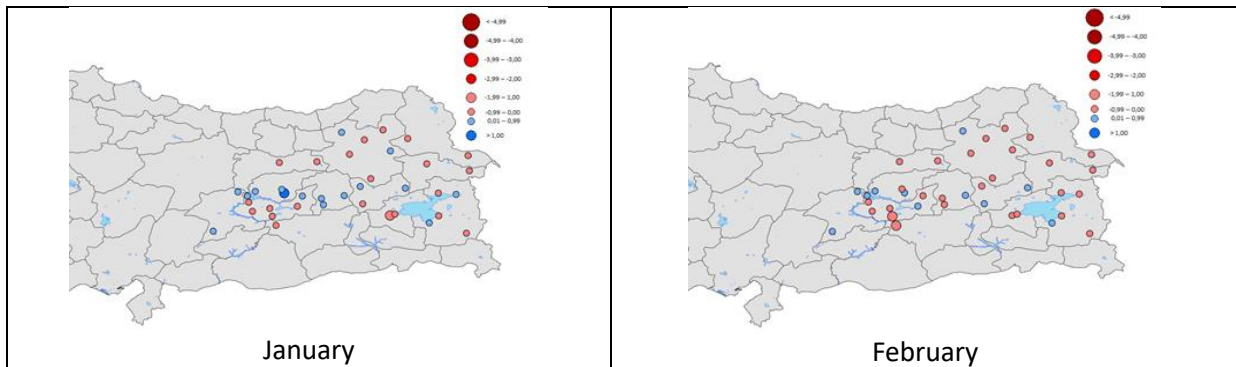
Figure 3.2. Spearman's rho test results for 10% and 5% level of significance

3.2.3. The slope of monthly total precipitation with Sen's slope method

The slopes of the monthly rainfall data at 38 rainfall measurement stations were estimated using Sen's slope method and the results are given in Table 3.3. The results are also shown in Figure 3.3.

Table 3.3. Sen's slope results

Station	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept.	Oct.	Nov.	Dec.
Ağrı	-0,34	-0,37	-0,27	0,32	0,30	-0,04	0,05	0,03	0,11	0,09	-0,17	-0,22
Doğubeyazıt	-0,02	-0,02	0,29	0,51	-0,13	-0,27	0,24	0,05	0,05	0,07	0,19	0,10
Erzincan	-0,06	-0,08	0,11	0,16	-0,28	-0,22	0,03	0,00	0,12	0,05	0,09	-0,21
Tercan	-0,16	0,00	-0,23	0,18	-0,53	-0,12	0,13	0,00	0,07	0,28	-0,25	-0,33
Erzurum	-0,12	-0,21	-0,06	0,32	0,12	-0,35	-0,10	-0,04	0,07	0,38	-0,14	0,03
Hınıs	-0,06	-0,07	-0,20	0,31	-0,19	-0,14	0,10	0,10	0,09	0,12	-0,29	-0,08
Horasan	0,07	-0,19	-0,02	0,42	-0,29	-0,15	0,33	0,17	0,22	-0,07	-0,23	-0,01
İspir	0,06	0,08	0,24	0,26	0,07	-0,11	0,43	0,03	-0,08	0,46	0,15	-0,42
Oltu	-0,14	-0,05	0,23	0,46	0,09	-0,18	0,40	0,05	0,12	-0,04	-0,20	-0,17
Tortum	-0,12	-0,16	-0,15	-0,05	-0,17	-0,38	0,30	-0,07	-0,07	-0,07	-0,26	-0,23
İğdir	-0,01	-0,08	0,06	0,28	0,24	-0,02	0,07	0,03	0,02	0,02	-0,04	-0,06
Sarıkamış	-0,07	-0,31	-0,17	0,21	-0,36	-0,53	0,16	-0,11	-0,07	0,14	-0,31	-0,32
Bingöl	0,00	-0,16	-0,32	-0,96	-0,17	-0,04	0,00	0,00	0,02	-0,15	-0,47	-0,34
Genç	0,81	-0,72	-1,28	-0,64	-0,95	-0,30	-0,03	0,00	0,08	-0,80	-2,16	-0,07
Solhan	0,25	0,42	0,35	-0,55	-0,09	0,11	0,01	0,00	0,08	-0,31	-0,87	-1,06
Elazığ	-0,23	-0,19	-0,64	-0,31	-0,22	-0,04	0,00	0,00	0,04	0,14	0,04	-0,33
Ağın	0,12	0,07	-0,07	-0,35	-0,49	-0,45	0,00	0,00	0,14	-0,43	-1,34	-0,01
Baskil	-0,29	-0,15	-0,54	-0,21	-0,62	-0,16	0,00	0,00	0,10	-0,22	-0,67	-0,15
Karakoçan	0,38	-0,46	-0,13	0,74	-1,56	-0,11	0,00	0,00	0,20	-0,19	-1,00	0,15
Keban	-0,20	-0,21	-0,46	-0,40	-0,49	-0,16	0,00	0,00	0,01	0,10	0,10	-0,15
Maden	-0,32	-1,80	-1,20	-0,87	-0,77	-0,16	0,00	0,00	0,06	0,72	-1,59	-1,31
Palu	-0,37	0,15	-0,22	-0,10	-0,46	-0,05	0,00	0,00	0,00	0,17	-0,63	-0,71
Sivrice	-0,26	-1,01	-0,63	-0,75	-0,90	-0,27	0,00	0,00	0,10	-0,24	-0,56	-1,24
Arapgir	0,11	0,05	-0,93	-0,67	-0,05	-0,10	0,00	0,00	0,07	0,32	-0,35	-0,69
Doğanşehir	0,14	0,48	-0,58	-0,64	-0,50	-0,07	0,00	0,00	0,04	-0,26	-0,36	-0,48
Tunceli	0,60	-0,11	-0,99	-0,59	-0,49	-0,10	0,00	0,00	0,13	-0,06	-1,19	-1,35
Çemişgezek	0,12	0,10	-0,78	-0,65	-0,06	-0,16	0,00	0,00	0,08	-0,18	0,24	-0,32
Mazgirt	1,68	0,53	-0,16	-0,32	-0,84	-0,41	0,00	0,00	0,13	-1,01	-2,00	0,46
Bitlis	-1,15	-0,66	0,27	-1,08	0,02	0,14	0,02	0,00	0,08	-0,83	-0,12	-0,94
Tatvan	-0,56	-0,48	-0,08	0,46	-0,07	-0,09	0,07	0,00	0,03	-0,63	-1,04	-0,64
Muş	-0,19	0,11	0,15	-0,30	-0,08	0,02	0,01	0,01	0,00	-0,36	-0,74	-0,54
Malazgirt	0,09	0,05	0,19	0,52	0,03	-0,17	-0,02	0,00	0,05	-0,32	-0,08	0,07
Varto	0,01	-0,36	0,48	0,09	-0,32	-0,25	0,28	0,20	0,33	0,01	-1,11	-0,15
Van	-0,01	-0,08	0,13	-0,27	0,09	-0,04	0,01	0,00	0,00	-0,10	-0,16	0,18
Başkale	-0,05	-0,07	-0,25	0,00	-0,23	-0,12	0,25	0,01	0,00	-0,47	0,12	-0,21
Erciş	-0,02	-0,15	-0,05	-0,17	-0,37	-0,46	0,12	0,00	0,02	-0,62	-0,57	-0,36
Gevaş	0,61	0,03	0,11	0,16	-1,19	-0,18	0,14	0,00	0,14	-0,96	-1,05	0,08
Muradiye	0,21	-0,02	0,42	0,23	0,07	0,11	0,19	0,00	0,12	-0,18	-0,03	-0,07



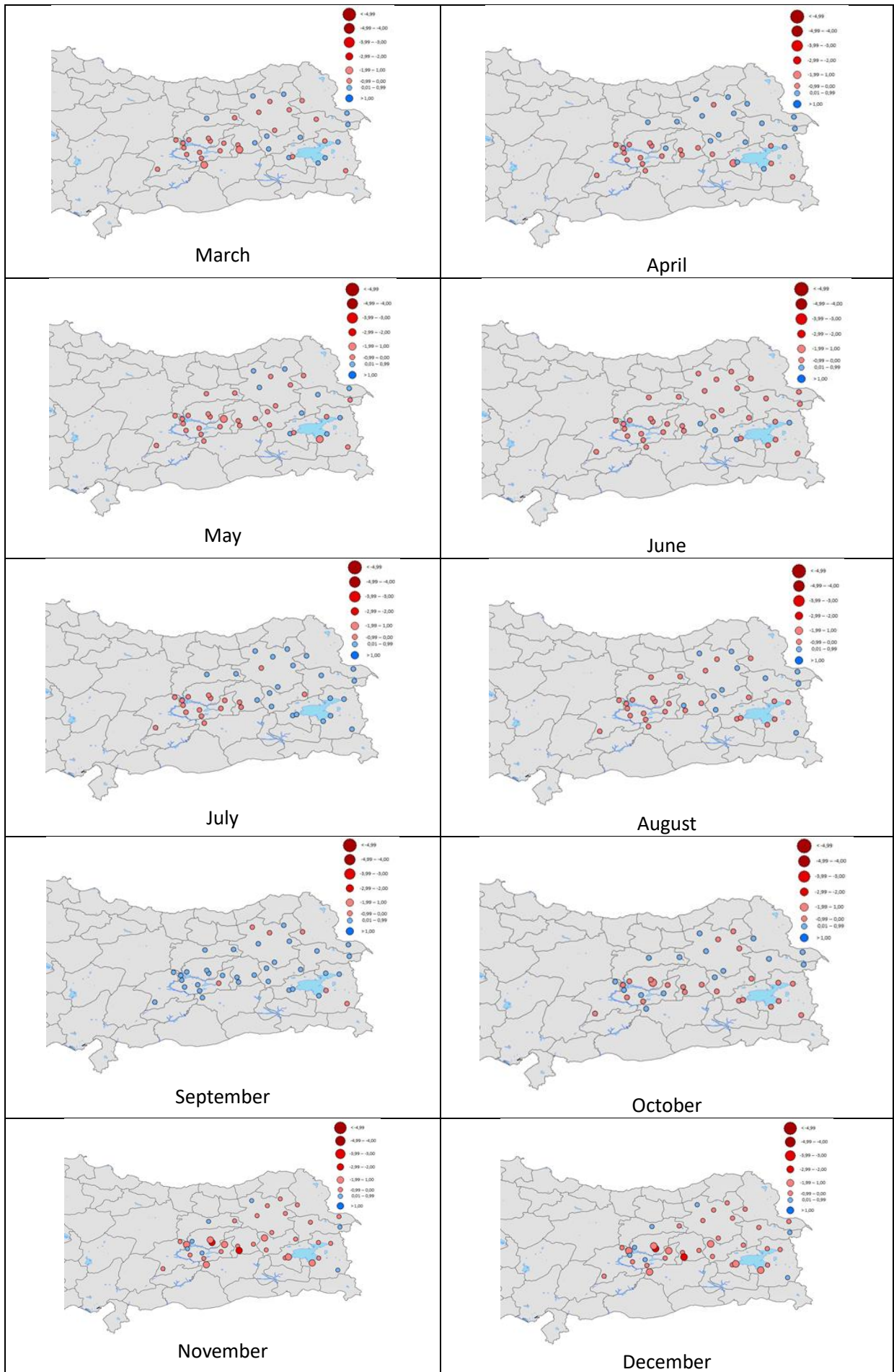


Figure 3.3. Monthly Sen's slope results

4. Conclusion

Rainfall is the most important environmental factor that determines the structural characteristics, species diversity and vital rhythm of the animal and plant communities on earth as it is the source of the water required for their living. Monitoring the expected or observed changes in precipitation regime and rainfall is especially important for the regions where the drought constitutes a threat.

In Eastern Anatolia, the largest geographical region of Turkey in surface area, precipitations are generally in the form of snow and uneven; therefore, the flow regime is also irregular. Since the precipitations are usually in snow form, in especially spring months, stream flows are rapid and strong. In addition, high elevations and slopes in the region also increase the hydroelectric potential.

In this study, the data obtained from the General Directorate of Meteorology were evaluated on monthly basis. As the result of the evaluation, it was observed that the largest rainfall was measured in April while the lowest was in August. The highest rainfall based on stations was observed in Bitlis station with 586.5 mm in February 1992 while the lowest (no rainfall) was measured in several stations in many months.

Two different methods were used to test the homogeneity of the data; Run test and Pettitt test. As a result of homogeneity analysis with Run test, it was determined that only Hakkari station was not homogeneous in 95% confidence interval. When the homogeneity analysis was performed by Pettitt test, it was determined that 7 rainfall stations, Ahlat, Ardahan, Arpaçay, Kars, Malatya, Özalp and Yüksekova, were not homogeneous in the 95% confidence interval.

Thus, among a total of 46 rainfall measurement stations, eight with non-homogeneous rainfall data were excluded from the study and 38 rainfall measurement stations were analyzed for trend analysis.

Mann – Kendall and Spearman's rho tests and Sen's inclination method were used for trend analysis of monthly total precipitation.

According to the common results of both of the trend analyses (at 90% confidence interval) which were conducted using Mann - Kendall test and Spearman's rho test and evaluated in the 90% and 95% confidence intervals, a significant decrease was observed at the rainfall measurement stations of Ağın in June and November, Ağrı in February, Arapgir in March, Elazığ in March, Erciş in June and October, Erzincan in June, Erzurum in February, Genç in November, Gevaş in May, Karakoçan in May, Keban in March, Mazgirt in June, Solhan in November and December and Tercan in May, while a significant increase was found for the stations of Tunceli in September and November, Ağın in September, Başkale in September, Çemişgezek in September, Doğubeyazıt in March and April, Elazığ in July, Erciş in July, Hınıs in August, İspir in July, Mazgirt in January, Muş in August, Oltu in April and July, Van in September, Varto in July, August and September.

In the 95% confidence interval, a significant increase was seen for the Rainfall Measurement Stations of Ağın in September, Başkale in July, Çemişgezek in September, Doğubeyazıt in March, İspir in July, Mazgirt in January, Oltu in April, Varto in September while significant decrease was seen in Elazığ (March), Erciş (June), Genç, (November) and Solhan in December.

It may be stated when evaluated at 95% confidence interval according to Sen's slope method that the results generally show similarities between the trends found in both Mann - Kendall test and Spearman's rho test; however, in July and August, the results were found to be closer to those found in the Mann - Kendall test method.

When the trend analysis was conducted at 38 rainfall measurement stations using the Mann Kendall test and the Spearman's rho test, in both trend methods similar and differing results were obtained for monthly rainfalls. For some stations and months, there is a rising trend while a decreasing trend can be seen according to the other method. In a similar way, it is possible to see a trend in a decreasing direction at a station while it has an increasing characteristic in another method.

Analysing these differences seen in the trend analysis of the monthly rainfall using different trend methods in addition to the Mann - Kendall test and Spearman's rho test may give better results in terms of the accuracy of the results.

It is seen when the trend analysis of the monthly rainfall is evaluated that in summer months, rainfall is generally in an increasing trend while in winter its trend is in a decreasing direction. In June (summer), there is a declining trend in the region while in July the trend turns out to be an increasing way. In November, the trend is in decreasing direction again in the region.

In the present study, trend analysis was performed using only precipitation data. Conducting similar studies for other meteorological parameters may serve as a basis for both the reliability of the study and the studies on climate change. Likewise, evaluation of the present results considering different climate scenarios for the the region may also contribute to the determination of the water potential in the future in the region.

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